

Study of mid-latitude ionospheric convection  
during super quiet, quiet and disturbed period  
with the SuperDARN Hokkaido radar

Yun Zou and Nozomu Nishitani  
(STEL, Nagoya University)

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# Possible factors for generating mid-latitude ionosphere convection

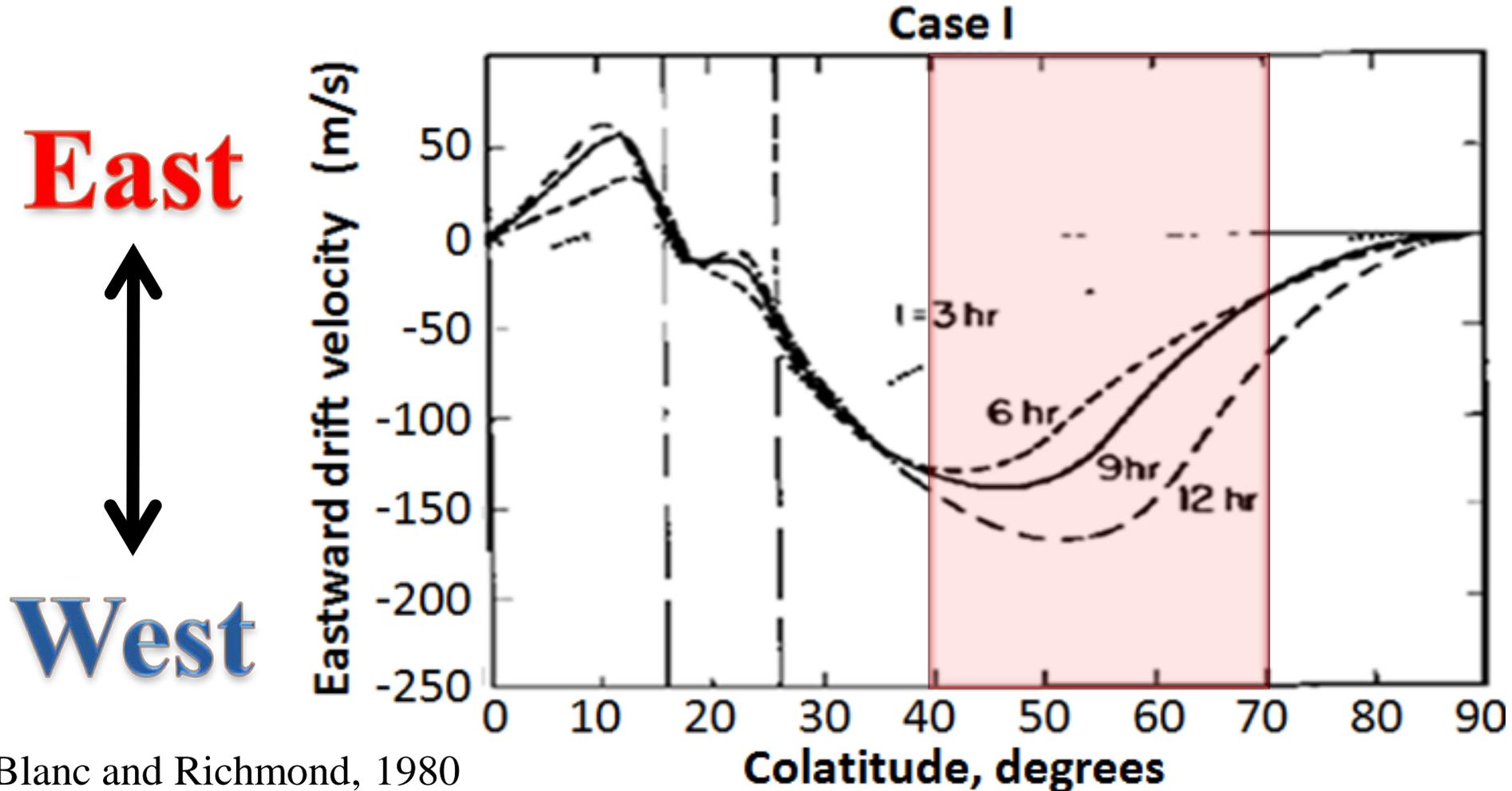
- Influence from high latitude convection (the penetration electric field)
- Formation of convection by tide (which dominate at low latitude)

- **Disturbance dynamo(which provides very large influence at mid-latitude)**

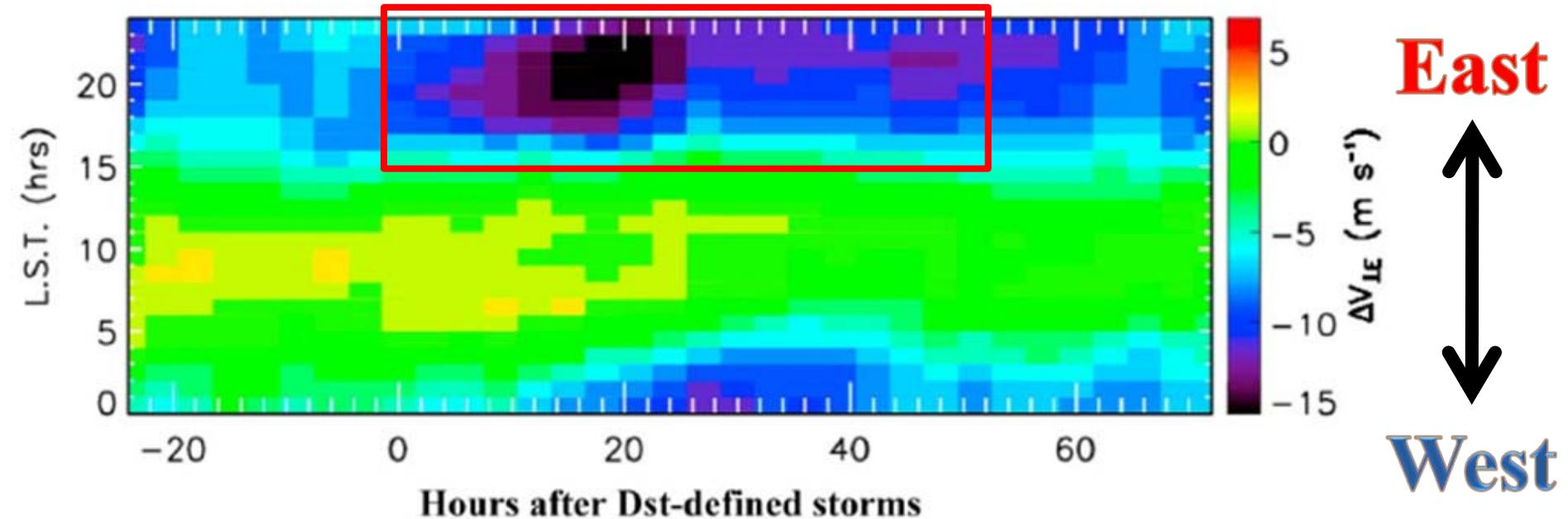
Our study of mid-latitude ionosphere convection will focus on Disturbance dynamo effects.

## Background(2/3)

- The result of the **simulation** shows that westward drift of F-region ionospheric plasma at mid-latitude develops with the growing of storms.
- In mid-latitude of Case I (**night side**), the speed of the **westward plasma drift** increases as the **geomagnetic disturbance develops**.



Ionospheric disturbance dynamos lasted up to **50 hours** in the midlatitude region at night

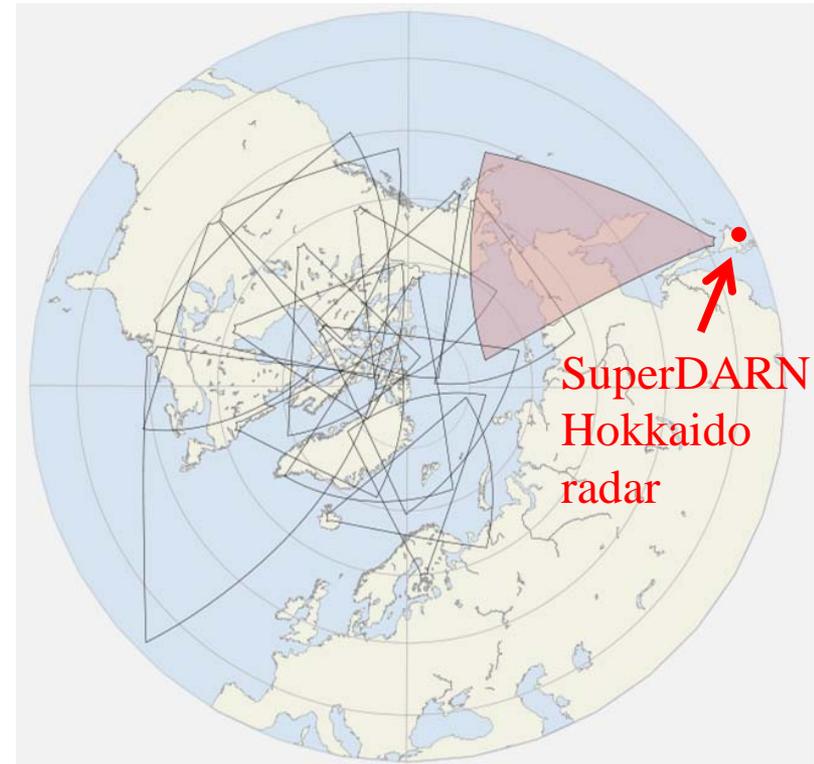


- Observation point : Bundoora ( $145.1^\circ$  E,  $37.7^\circ$  S geographic,  $49^\circ$  S magnetic), Australia
- Observation equipment : Digisonde drift measurements
- Statistical methods : Superposed Epoch Analysis (SEA)
- Database : 5 years (1999 to 2003)
- Number of Dst-defined storms : 262

# Purpose of this study

To study mid-latitude ionosphere plasma convection (especially the westward flow on the night side)

- K<sub>p</sub> dependence
- Geomagnetic latitude dependence
- Seasonal dependence
- Influence from Dst-defined storms
- Influence from AL-defined storms



## SuperDARN Hokkaido radar

➤ **Location** :  $43.5^{\circ}$  N geographic,  $36^{\circ}$  N magnetic

➤ **Observation area** : about  $40^{\circ}$  to  $80^{\circ}$  N

# Method of this study

To study the geomagnetic activity dependence and the geomagnetic latitude dependence of the mid-latitude ionospheric convection, we apply the **vector fitting** to **beam 1 to 14** line-of-sight Doppler velocity data obtained by the SuperDARN Hokkaido radar to get the **westward and northward components** of ionospheric convection velocities.

- In order to minimize the effects of sidelobes, data from beam 0 and 15 are not used.
- For beam 1 to 14, the vector can be obtained for eastward and northward components from the equations:

$$v_1 = A \cos \theta_1 + B \sin \theta_1$$

$$v_2 = A \cos \theta_2 + B \sin \theta_2$$

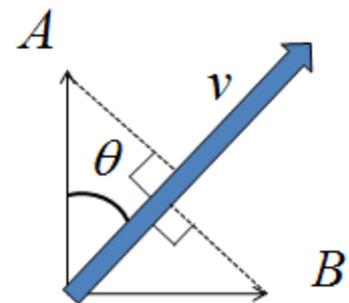
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$$v_{14} = A \cos \theta_{14} + B \sin \theta_{14}$$

- To Calculate these vector components we use equation:

$$B = \frac{(\sum v_i \cos \theta_i)(\sum \cos \theta_i \sin \theta_i) - (\sum v_i \sin \theta_i)(\sum \cos^2 \theta_i)}{(\sum \cos \theta_i \sin \theta_i)^2 - (\sum \sin^2 \theta_i)(\sum \cos^2 \theta_i)}$$

$$A = \frac{(\sum v_i \sin \theta_i)(\sum \cos \theta_i \sin \theta_i) - (\sum v_i \cos \theta_i)(\sum \sin^2 \theta_i)}{(\sum \cos \theta_i \sin \theta_i)^2 - (\sum \sin^2 \theta_i)(\sum \cos^2 \theta_i)}$$



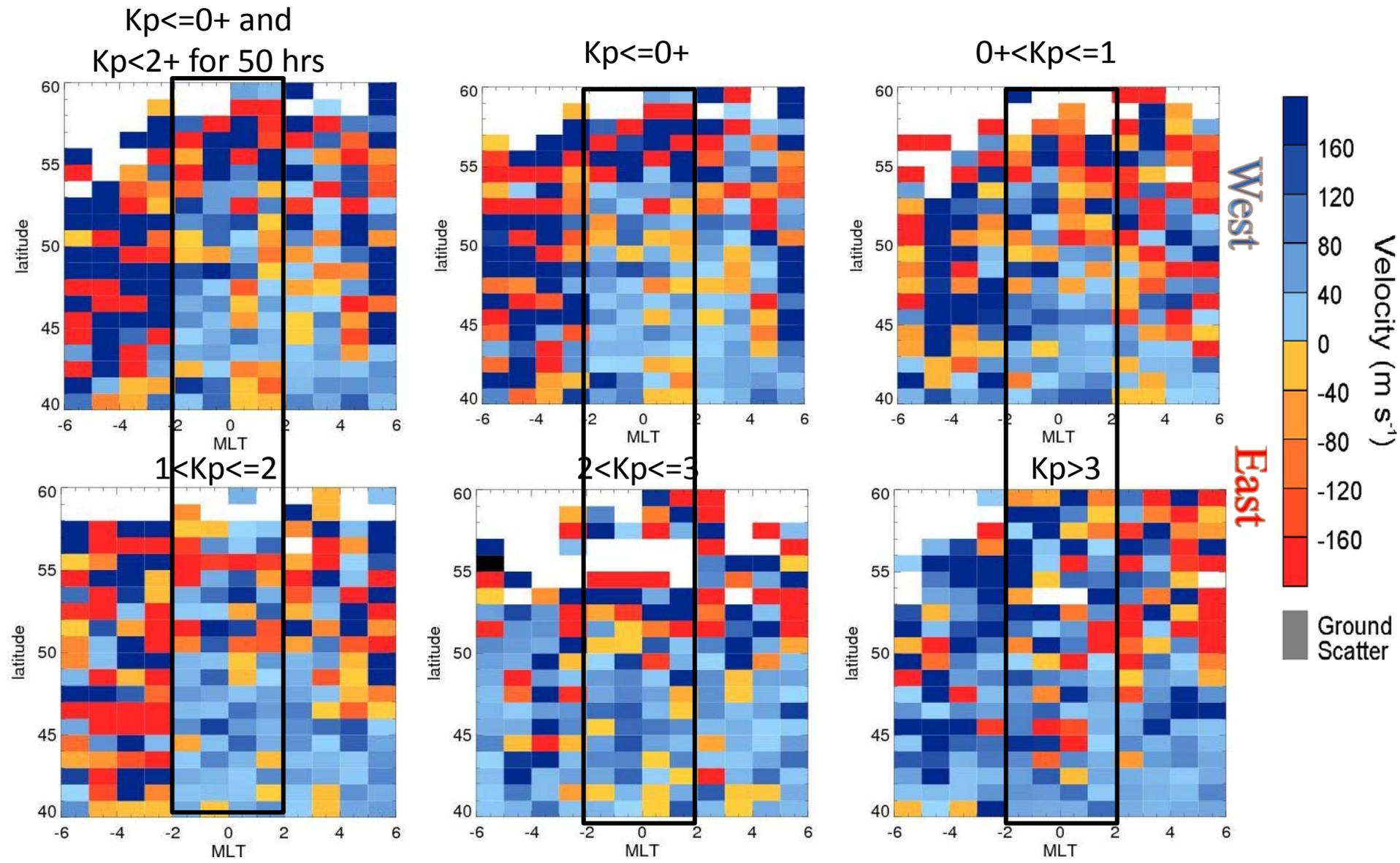
Here,  $\theta$  is the azimuth at each range of each beam.

# Result

- Kp dependence
- Geomagnetic latitude dependence
- Seasonal dependence
- Influence from Dst-defined storms
- Influence from AL-defined storms

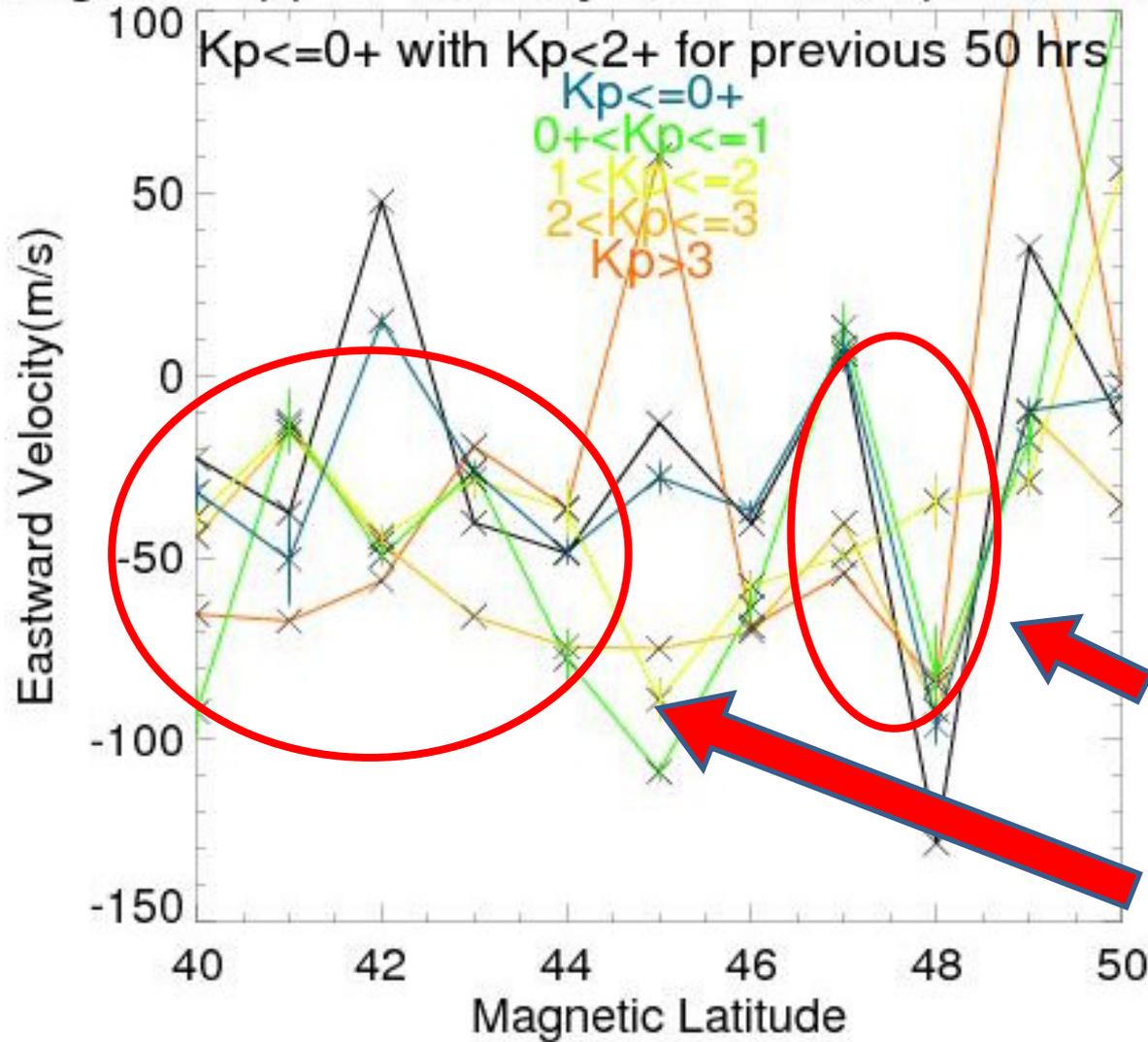
# Result(1/4)

Statistical distribution of **east-west** velocities with several Kp ranges in **winter** (Nov to Jan) from December, 2006 to May, 2011.



## Result(2/4)

Averaged Doppler velocity with each kp for 22 to 02 MLT



**East**



**West**

With increasing Kp index, **westward** flow becomes more intense in the nightside mid-latitude region.

Because the data count of the averaged Doppler velocity for Kp>3 is much less than other conditions, so that the variation is much greater.

# Result

- Kp dependence
- Geomagnetic latitude dependence
- Seasonal dependence
- Influence from Dst-defined storms
- Influence from AL-defined storms

# Criteria

● To study the geomagnetic latitude dependence of the chronological influence of **magnetic storms** and **substorms**, we use the **Superposed Epoch Analysis (SEA)** method .

Storm Time(t)	All Storms	Short-Duration	Medium-Duration	Long-Duration
$-7h \leq t \leq -1h$	$Dst > -30nT^{*1}$	$Dst > -30nT^{*1}$	$Dst > -30nT^{*1}$	$Dst > -30nT^{*1}$
$0h \leq t \leq 6h$	$Dst \leq -30nT^{*2}$	$Dst \leq -30nT^{*2}$	$Dst \leq -30nT^{*2}$	$Dst \leq -30nT^{*2}$
$7h \leq t \leq 13h$	—	$Dst > -60^{*3}$	$Dst \leq -60^{*4}$	$Dst \leq -60^{*4}$
$14h \leq t \leq 20h$	—	—	$Dst > -60^{*3}$	$Dst \leq -60^{*4}$
Number of Storms	54	48	4	2

\*1 : 1 or less in seven hours satisfy the stated condition

\*2 : 3 or less in seven hours satisfy the stated condition

\*3 : All 7 consecutive hours satisfy the stated condition

\*4 : At lest 1 hour satisfy the stated condition

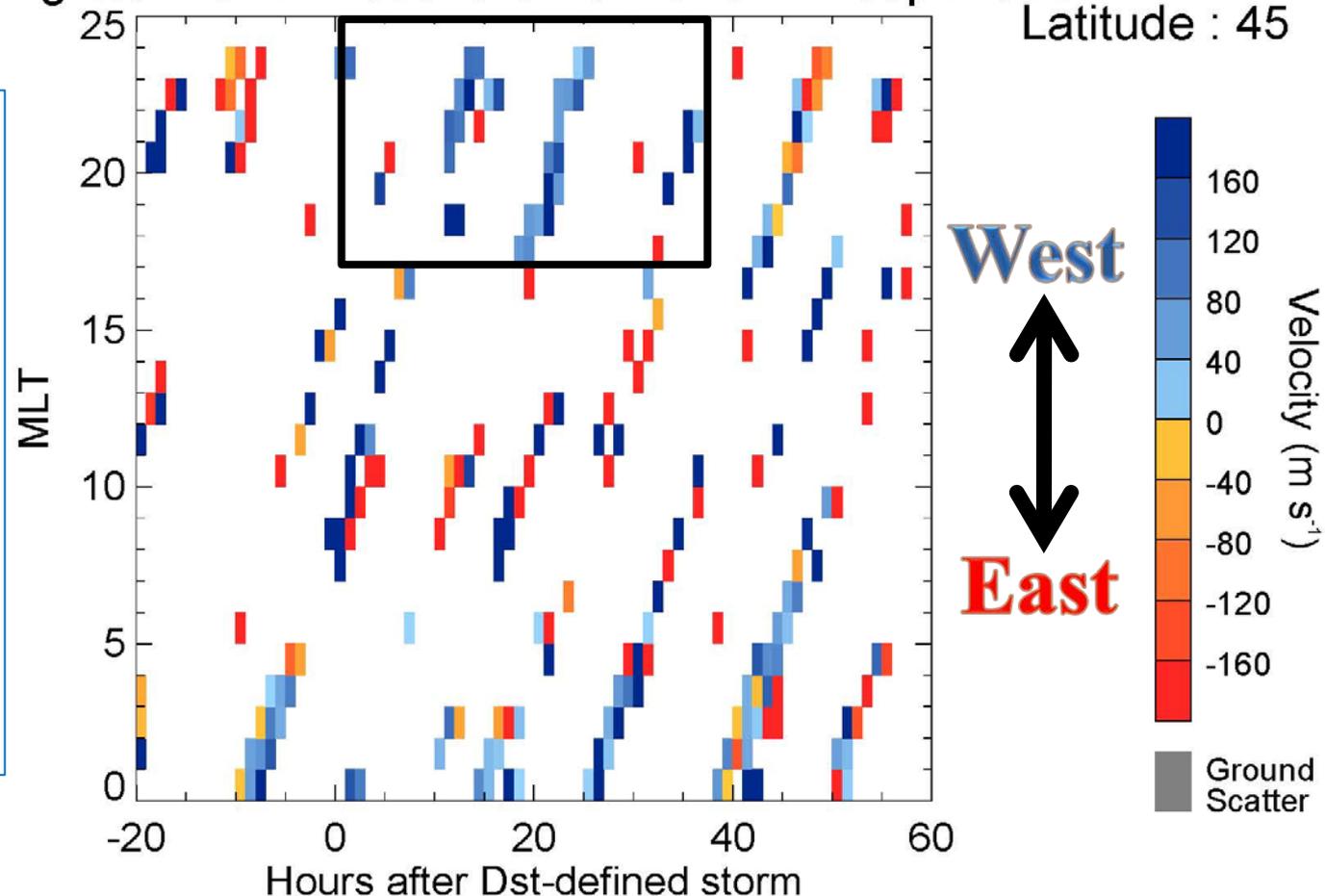
Storm Time(t)	$-6h \leq t \leq 0h$	$0h < t \leq 6h$
All substorms	$AL > -150$	$AL \leq -150$

Number of Storms : 94

## Result(3/4)

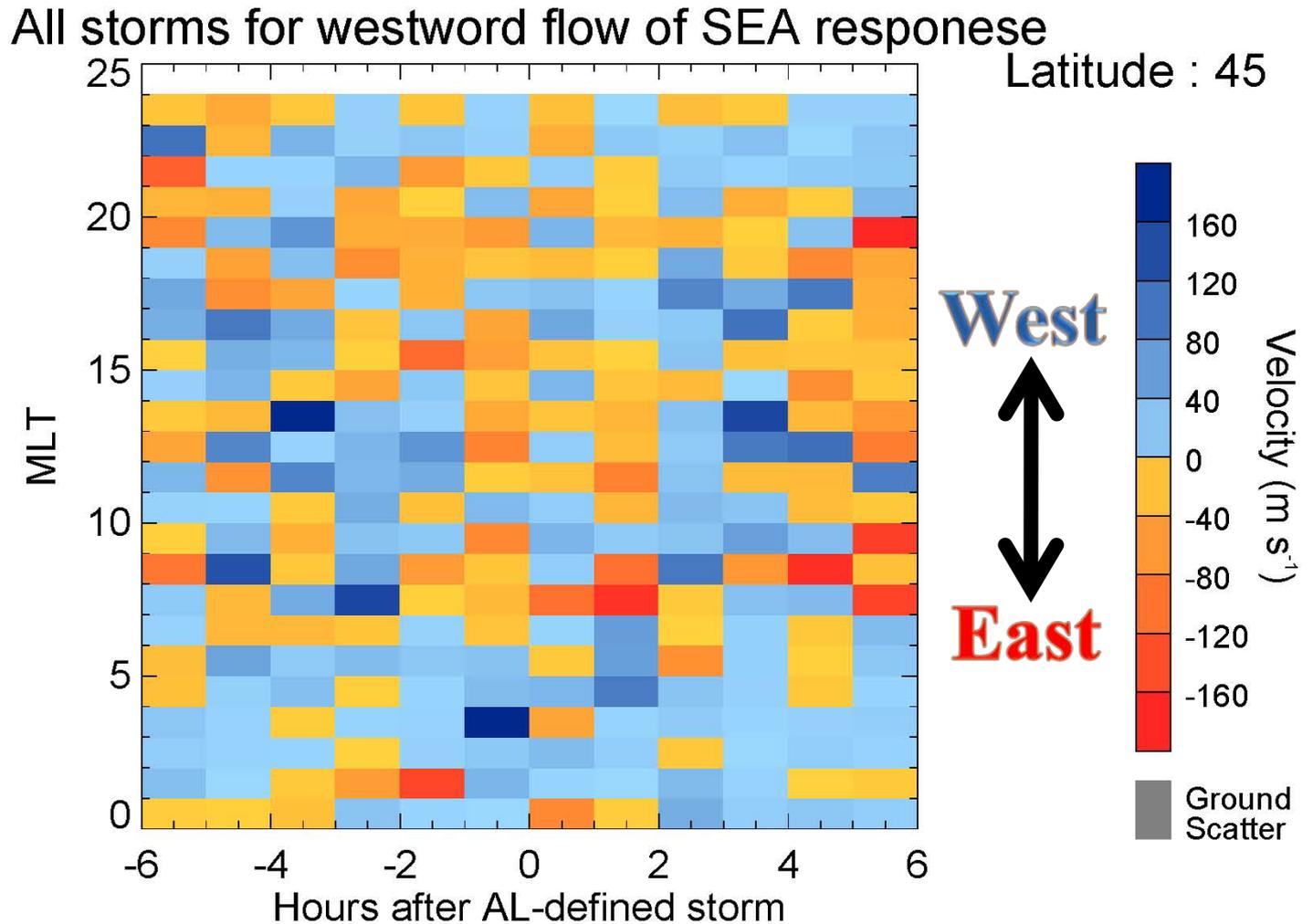
From the result of 6 events of medium- and long-duration storms we can confirm that the influence of storm **lasts up to 30 hrs** after onset between **43** and **59** degrees.

Big storms for westward flow of SEA response



- 54 storms defined by the hourly Dst index which satisfy the criteria, from December, 2006 to May, 2011.
- We use the criteria adopted by Kumar et al., 2010.

Statistical analysis of 94 AE-defined substorms 6 hours before and after the substorm onset (12 hours total)



The result was not particularly noticeable.

There is a need for reconsider on a detailed understanding of the features of substorm and it's influence on mid-latitude ionospheric convection.

# Summary

- The statistical analysis shows the **geomagnetic index Kp and geomagnetic latitude dependence** of the mid-latitude ionospheric plasma convection from the observation of SuperDARN Hokkaido radar **in all seasons**. The mid-latitude **night side westward flow** has **Kp dependence** intensified for disturbed periods in all seasons.
- The **Superposed Epoch Analysis (SEA) result** shows a **latitude dependence** of the influence from the storm. From the result of 6 events of medium- and long-duration storms we can confirm that the influence of storm **lasts up to 30 hrs** after onset at several latitude of mid-latitude.
- Significant influence from the AL-defined substorm at mid-latitude is not identified in **6 hours after substorm begins**. Kumar et al. [2010] shows the most great effect from the AE-defined storm is at **20 hrs after the substorm onset**. The simulation result from Blanc and Richmond [1980] also shows the westward drift at mid- and low-latitude grows to a maximum at **12 hrs after the magnetic storm onset**. There is a need for reconsider on a detailed understanding of the features of substorm and its influence on mid-latitude ionospheric convection.

## Future work

- **More events** of each duration storms need to be studied to extend the detailed result of data analysis.
- **Other possible factors** to affect mid-latitude convection besides the disturbance dynamo should be investigated.





# Overview of disturbance dynamo at mid-latitude from Previous studies

- Dominates at mid-latitude
- Provides strong influence of westward flow
- The influence becomes larger with the growth of storms
- The influence may last for 50 hours

Geomagnetic latitude dependence was not shown by the observation

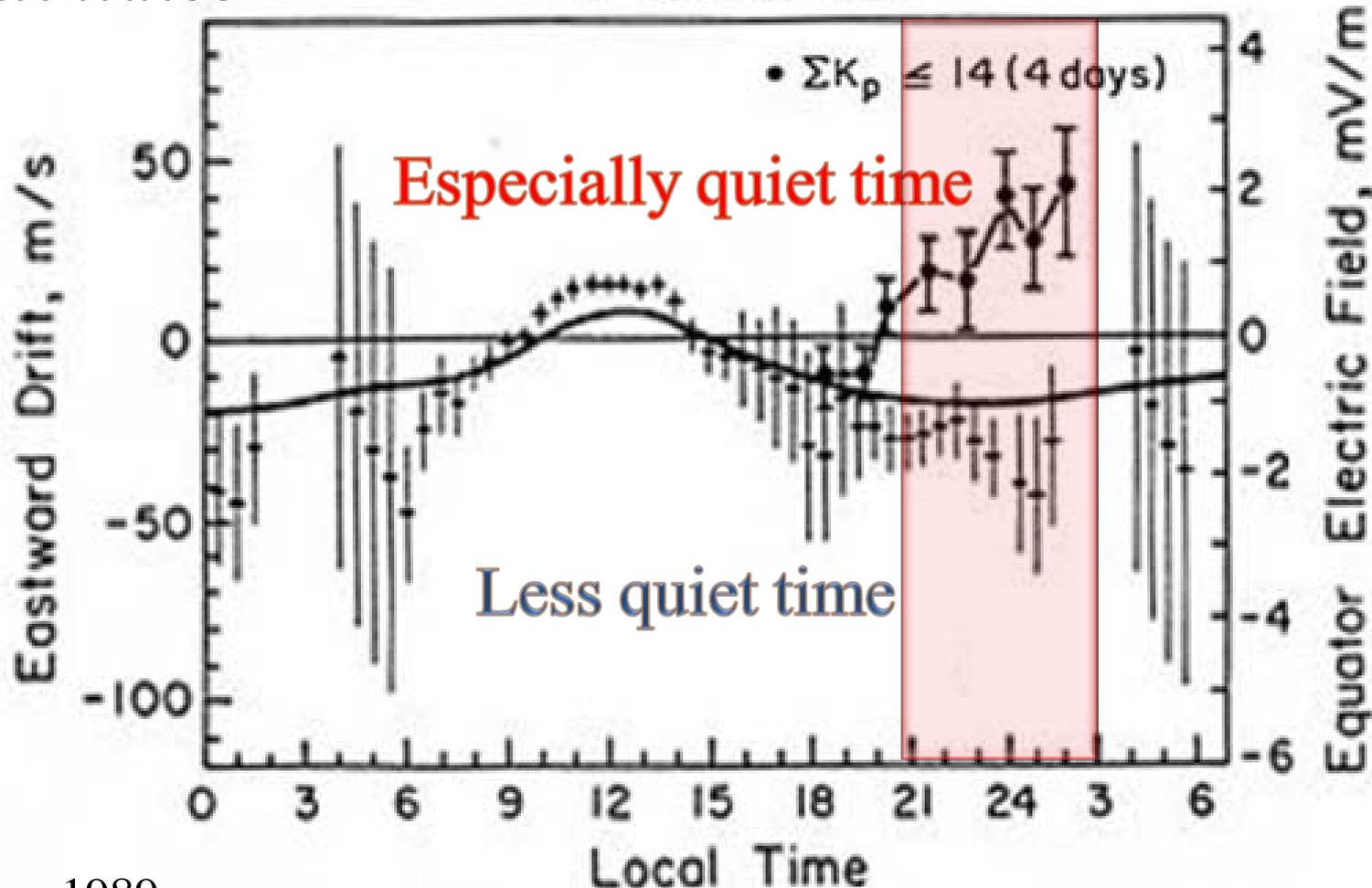
# Background(4/4)

Observation result using Millstone Hill radar by **Gonzales et al. [1978]** and **Richmond et al. [1980]**

Geomagnetic latitude:57

Millstone Hill And the empirical model

**East**  
↑  
↓  
**West**



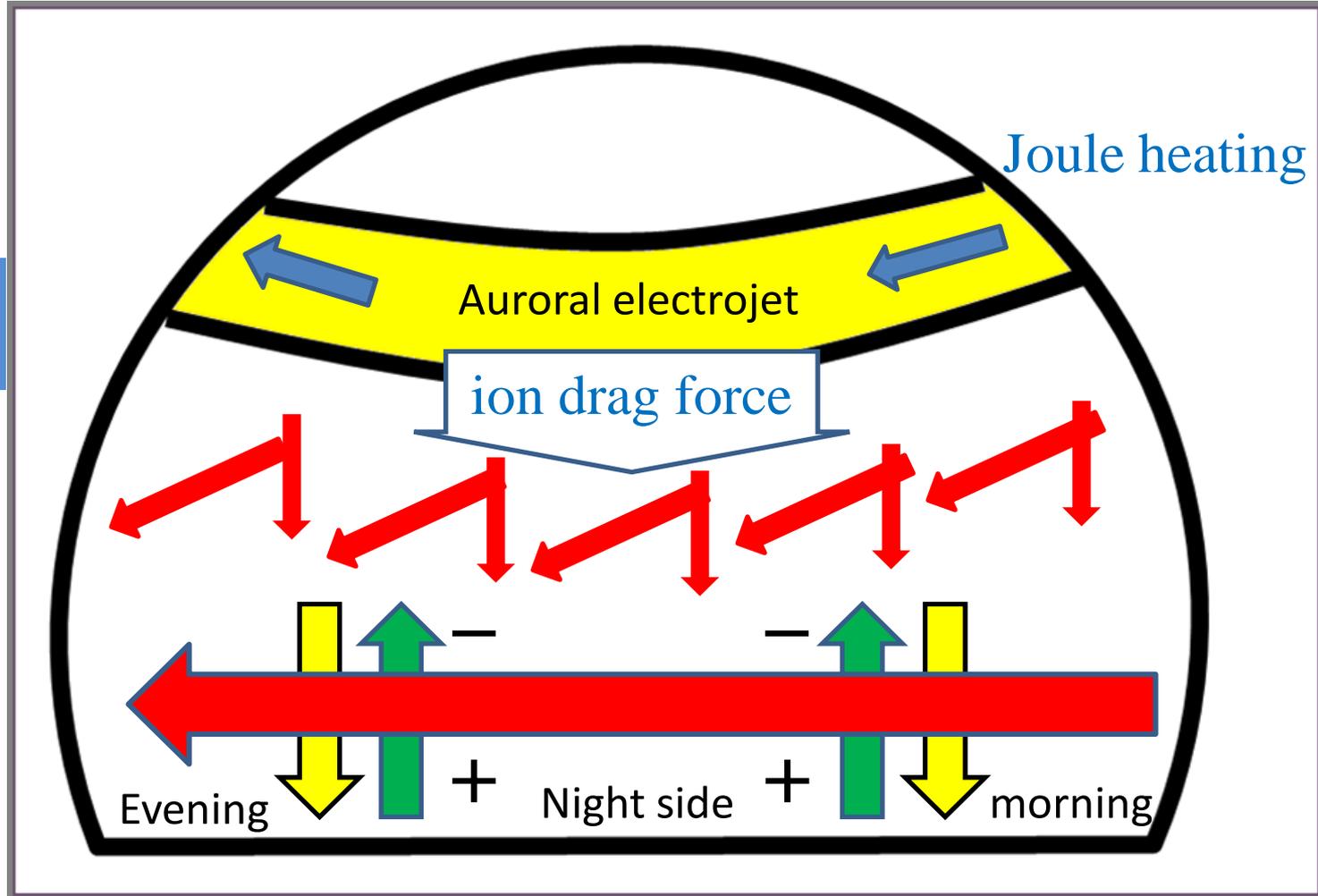
# The Mechanism of disturbance dynamo

Coriolis force  
change southward  
neutral wind to  
westward

$U \times B$  electric field  
(Pedersen current)

Charge appears  
due to the  
conductivity  
gradient

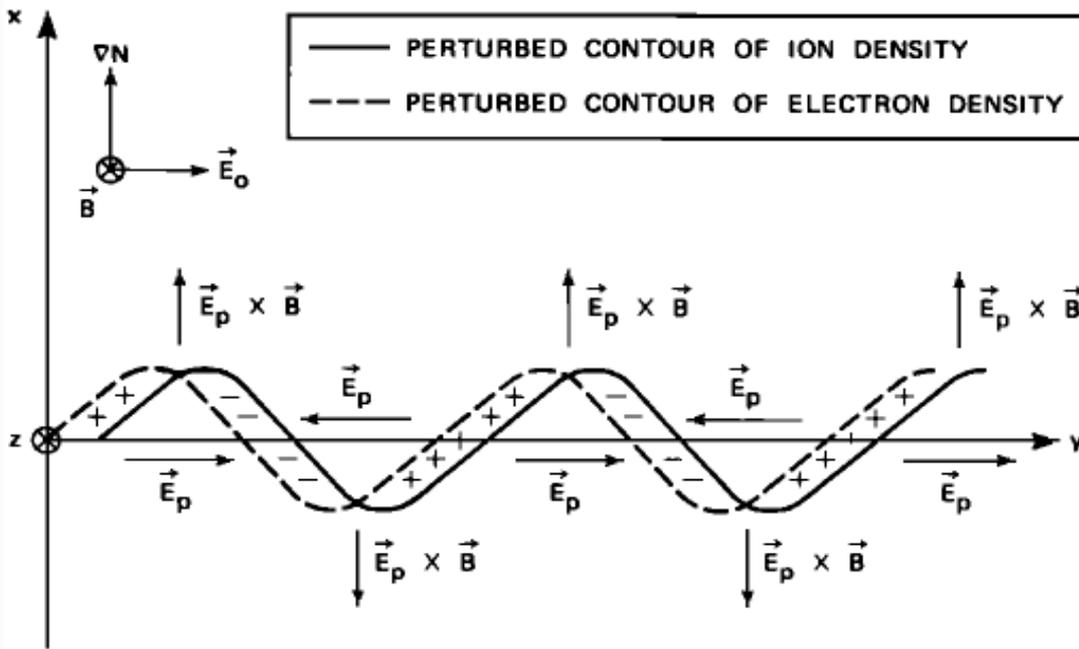
westward  $E \times B$   
drift in the mid-  
latitude



# Discussion

## Comparison with past studies

- ◆ In contrast to the result of Gonzales et al., night side eastward flow was not observed during deadly quiet periods.



A Pedersen ion drift (to the right) leads to charge separation and the development of polarization electric fields,  $E_p$ . The sense of  $E_p$  drives  $E_p \times B$  motion that further enhances the original plasma perturbation.

TSUNODA, 1988

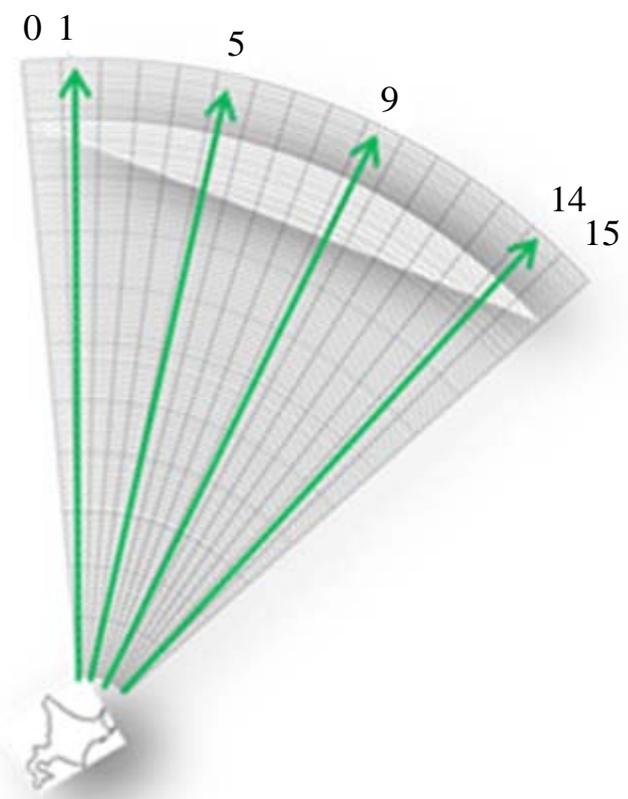
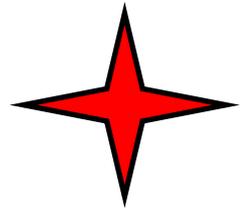
## Why?

- One important factor of **gradient drift instability** is most likely to generate **ionospheric irregularities**.
- Gradient drift instability occurs when the **electric field is strong enough**, so that the ionospheric echoes are observed preferably when the electric field is strong (Fukumoto et al., 2000).
- **Westward flow** is observed preferably when the electric field is strong.

# Method of this study(2/3)

Here  $\theta$  means the zenith angle of each beam at different range, here we are using Spherical trigonometry to calculate the zenith angle.

Magnetic north



SuperDARN Hokkaido radar

Radar beam angle height=110km

